Batch extraction method to estimate total dissolved solids (TDS) release from coal refuse and overburden[☆]

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ABSTRACT

A rapid batch extraction method was evaluated to estimate potential for total dissolved solids (TDS) release by 65 samples of rock from coal and gas-bearing strata of the Appalachian Basin in eastern USA. Three different extractant solutions were considered: deionized water (DI), DI equilibrated with 10% CO_2 atmosphere (DI + CO_2), or 30% H_2O_2 under 10% CO_2 (H_2O_2+CO_2). In all extractions, 10 g of pulverized rock (<0.5-mm) were mixed with 20 mL of extractant solution and shaken for 4 h at 50 rpm and 20–22 °C. The 65 rock samples were classified as coal (n=3), overburden (n = 17), coal refuse that had weathered in the field (n = 14), unleached coal refuse that had oxidized during indoor storage (n = 20), gas-bearing shale (n = 10), and pyrite (n = 1). Extracts were analyzed for specific conductance (SC), TDS, pH, and major and trace elements, and subsequently speciated to determine ionic contributions to SC. The pH of extractant blanks decreased in the order DI (6.0), DI + CO_2 (5.1), and H_2O_2+CO_2 (2.6). The DI extractant was effective for mobilizing soluble SO_4 and Cl salts. The DI + CO_2 extractant increased weathering of carbonates and resulted in equivalent or greater TDS than the DI leach of the same material. The H_2O_2+CO_2 extractant increased weathering of sulfides (and carbonates) and resulted in the greatest TDS production and lowest pH values. Of the 65 samples, 19 had leachate chemistry data from previous column experiments and 35 were paired to 10 field sites with leachate chemistry data. When accounting for the water-to-rock ratio, TDS from DI and DI + CO_2 extractions were correlated to TDS from column experiments while TDS from H_2O_2+CO_2 extractions was not. In contrast to column experiments, field SC was better correlated to SC measured from H_2O_2+CO_2 extractions than DI extractions. The field SC and SC from H_2O_2+CO_2 extractions were statistically indistinguishable for 7 of 9 paired data sets while SC from DI extractions underestimated field SC in 5 of 9 cases. Upscaling comparisons suggest that (1) weathering reactions in the field are more aggressive than DI water or synthetic rainwater extractants used in batch or column tests, and (2) a batch extraction method utilizing 30% H_2O_2 (which is mildly acidic without CO_2 enrichment) could be effective for identifying rocks that will release high amounts of TDS.

1. Introduction

High salinity in streams downgradient of coal-mining and processing facilities in the eastern United States (U.S.) has caused fish kills and harmed sensitive aquatic organisms within the past decade (Barrett, 2015; Cormier et al., 2013a,b; Pond et al., 2008). At the same time,

Table 1
Descriptions of the 65 rock samples tested.

ID	Source[a]	Operational Rock Type[b]	Lithology	WE[c]	Stratigraphic Formation	Local Geologic Description	Mineralogy[d]	Data for Upscaling[e]
PA1	Mine A	Coal	Coal	W	Kittanning	L. Kittanning	Qtz, Cal, Kln, Py, Ms	n.a. [f]
PA5	Mine A	Coal	Coal	W	Kittanning	L. Kittanning	Qtz, Kln, Py, Ms	n.a.
PA45	Mine B	Coal	Coal	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Kln, Jr, Ms	n.a.
PA12	Mine A	W. Refuse	Coal and shale	W	Kittanning	L. Kittanning	Qtz, Gp, Kln, Ms, Jr,	FL
PA13	Mine A	W. Refuse	Coal and shale	W	Kittanning	L. Kittanning	Qtz, Gp, Kln, Ms, Jr,	FL
PA17	Mine A	W. Refuse	Coal and shale	W	Kittanning	L. Kittanning	Qtz, Kln, Ms, Py,	FL
PA22	Mine A	W. Refuse	Coal and shale	W	Kittanning	L. Kittanning	Qtz, Ms, Kln, Py, Gp, Vrm	FL
PA30	Mine A	W. Refuse	Coal and shale	W	Kittanning	L. Kittanning	Qtz, Cal, Kln, Py, Ms, Gp	FL
PA31	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Gp, Ms, Kln, Py, Cal, Vrm	FL
PA36	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Kln, Ms, Py, Vrm	FL
PA42	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Klm, Ms, Gp, Vrm	FL
PA48	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Ms, Kln, Gp, Py, Vrm	FL
PA51	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Ms, Kln, Gp, Vrm	FL
PA58	Mine B	W. Refuse	Coal and shale	W	Kittanning and Allegheny	L. Kittanning/U. Kittanning/Freeport	Qtz, Ms, Kln, Gp, Jr, Vrm	FL
TNR1	VT	W. Refuse	n.a.	n.a.	Graves Gap	Graves Gap refuse	Qtz, Ms, Kln, Vrm	CL
TNR2	VT	W. Refuse	n.a.	n.a.	Graves Gap	Graves Gap refuse	Qtz, Cal, Ms, Kln, Gp, Vrm	CL
TNR3	VT	W. Refuse	n.a.	n.a.	Graves Gap	Graves Gap refuse	Qtz, Ms, Kln, Vrm, Py	CL
TGS1	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Py, Kln, Ms, Vrm	FL
TGS2A	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Poi, Py, Kln	n.a.
TGS2B	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms	FL
TGS3	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Cal, Kln, Ms, Py	FL
TGS4	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms, Vrm	n.a.
TGS5	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Poi, Py, Kln	FL
TGS6	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms	FL
TGS7A	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Cal, Ms, Kln	FL
TGS7B	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Kln, Py, Ms	FL
TGS8	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms,Poi, Rz	FL
TGS9	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms	FL
TGS10A	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Rz, Kln, Py, Ms	n.a.
TGS10B	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms, Rz	FL
TGS11	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Qtz, Kln, Ms	FL
TGS12	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Qtz, Gp, Py, Kln, Ms	FL
TGS13	PA	U. Refuse	Shale	U	Monongahela	Pittsburgh	Cal, Qtz, Kln	FL

ID	SM[a]	Operational Rock Type[b]	Lithology	WE[c]	Stratigraphic Formation	Local Geologic Description	Mineralogy[d]	Data for Upscaling[e]
TGS14	PA	U. Refuse	Coal and shale	U	Monongahela	Pittsburgh	Qtz, Ms, Klm, Py, Vrm	FL
TGS15	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Klm, Ms, Py	FL
TGS16	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Py, Klm, Ms	FL
TGS17	PA	U. Refuse	Claystone	U	Monongahela	Pittsburgh	Qtz, Klm, Ms, Cal, Py	FL
BCS3	USGS	Overburden	Shale	U	Glenshaw	Brush Creek shale	Qtz, Ms, Kln, Chl, Vrm, Cal, Py	CL, FL
HCS	USGS	Overburden	Shale	U	Carbondale	Houchin Creek shale	Qtz, Gp, Ms, Ill, Chl, Kln, Jr, Py, Ab, Vrm	CL
KBFWV	USGS	Overburden	Shale	U	Kanawha	Black Flint shale	Qtz, Kln, Ms, Dol, Sd, Ab, Vrm	CL
LKFC	USGS	Overburden	Shale	U	Kittanning	L. Kittanning	Qtz, Ms, Chl, Ill, Kln, Vrm, Gp, Py, Sd	CL, FL
MKSS	USGS	Overburden	Sandstone	U	Kittanning	M. Kittanning	Qtz, Ms, Kln, Chl, Vrm, Gp, Cal, Sd	CL
KY1	UK	Overburden	Sandstone	W	Princess	Princess	Qtz, Kln, Vrm, Dol, Sd, Gt	CL, FLY
KY2	UK	Overburden	Sandstone	U	Four Corners	Four Corners	Qtz, Kln, Vrm, Dol, Sd	CL, FLY
KY3	UK	Overburden	Mixed	W	Four Corners	Four Corners	Qtz, Kln, Vrm, Dol, Sd, Gt	CL, FLY
KY4	UK	Overburden	Mixed?	U	Four Corners	Four Corners	Qtz, Kln, Vrm	CL

Table 1 (*continued*)

ID	Source[a]	Operational Rock Type[b]	Lithology	WE[c]	Stratigraphic Formation	Local Geologic Description	Mineralogy[d]	Data for Upscaling[e]
KY7	VT	Overburden	Mixed	U	Four Corners	Four Corners black shale	Qtz, Ms, Kln, Ab, Gp, Vrm	CL, FL
KY9	VT	Overburden	Mixed	U	Four Corners	Four Corners mixed	Qtz, Ms, Kln, Vrm	CL
TN2	VT	Overburden	Shale mix	n.a.	Anderson and Glen Dean	Windrock, Lower Dean, Dean	n.a.	CL
VA2	VT	Overburden	Black shale	U	Four Corners	Four Corners black shale	Qtz, Ms, Kln, Vrm	CL
VA3	VT	Overburden	Mixed	U	Wise	M. Wise mixed	Qtz, Ms, Kln, Ab, Vrm	CL
VA6	VT	Overburden	Mudstone	U	Wise	Lower Wise mudstones	Qtz, Ms, Kln, Vrm	n.a.
VA16	VT	Overburden	Sandstone	U	Harlan	Harlan Sandstone	Qtz, Ms, Kln, Vrm	CL
WV5	VT	Overburden	Sandstone	U	Kanawha	Kanawha Sandstone	Qtz, Ms, Kln, Vrm	CL
SHJ1	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Dol, Py, Vrm	n.a.
SHJ2	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Dol, Py, Vrm	n.a.
SHJ3	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Dol, Py, Vrm	n.a.
SHJ4	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Dol, Py, Vrm	n.a.
SHJ5	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Dol, Py, Vrm	n.a.
SHM1	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Kln, Ms, Dol, Py	n.a.
SHM2	NDA	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Kln, Ms, Dol, Py	n.a.
SHM3	NDA	Shale	Shale gas	U	Utica	Utica Shale	n.a.	n.a.
SHO1	OH	Shale	Shale gas	U	Utica	Utica Shale	Qtz, Cal, Ms, Kln, Vrm, Py, Dol	n.a.
SHE1	PA	Shale	Shale gas	U	Marcellus	Marcellus Shale	Qtz, Cal, Ms, Kln, Vrm, Gp, Py	n.a.
SKYPA	USGS	Pyrite	Pyrite	n.a.	Bald Eagle	Bald Eagle Sandstone	Qtz, Ms, Py	FR

[a] Source: OH = Ohio Geologic Survey; NDA = Penn State Non-disclosure agreement; PA = Pennsylvania Bureau of Topographic and Geologic Survey; USGS = U.S. Geological Survey; UK = University of Kentucky; VT = Virginia Tech.

[b] U. Refuse = Unleached refuse; W. Refuse = Weathered refuse.

[c] WE = Weathering extent: U = Unweathered (partly oxidized while stored indoors but unleached); W = Weathered (partly oxidized and leached outdoors).

[d] Minerals identified by XRD and are listed in semi-quantitative order of abundance. Ab = albite; Cal = calcite; Chl = chlorite; Dol = dolomite; Gp = gypsum; Gt = goethite; Ill = illite; Jr = jarosite; Kln = Kaolinite; Ms = muscovite; Poi = poitevinite; Py = pyrite; Qtz = quartz; Rz = rozenite; Sd = siderite; Vrm = vermiculite.

[e] CL = Column leachate; FL = Field leachate; FLY = Field lysimeter; FR = Field runoff.

[f] n.a. = not available.

higher than normal concentrations of total dissolved solids (TDS), chloride, and bromide have been documented in the Allegheny and Monongahela Rivers in western Pennsylvania (Wang, 2014; Ziemkiewicz, 2015a), while a gradual increase in salinity attributed to chloride has been documented in major rivers in the northeastern U.S. (Kaushal et al., 2005, 2018). Such observations for coal-mine drainage in the northern Appalachian Basin may be explained by recent changes in resource extraction activities that can be influenced by residual brine in the rock, including the underground mining of coal into progressively deeper zones, the development of coal-bed methane, and the development of shale gas reserves in strata below the coal-bearing formations, notably the Marcellus Shale (Cravotta and Brady, 2015; Donovan and Leavitt, 2004; Donovan et al., 2015; Ziemkiewicz, 2015a). An understanding of the potential sources of salinity at local and watershed scales is necessary for the development of effective strategies to minimize and mitigate aquatic impacts from elevated TDS.

Accelerated mineral weathering generally accounts for increased TDS release from coal-mining landscapes (Brady et al., 1998; Timpano et al., 2010, 2015). Although acidic drainage and TDS release are commonly attributed to abandoned mines, the use of overburden materials as topsoil substitutes and the placement of carbonate-bearing overburden materials can contribute to elevated TDS (Bernhardt et al., 2012; Cormier et al., 2013a,b; Zipper et al., 2015). Topsoil substitution with overburden is explicitly allowed in the Surface Mining Control and Reclamation Act (SMCRA), and the application of acid-base accounting (ABA) procedures guides placement of alkaline strata with the explicit goal of decreasing acidity from sulfide oxidation (Skousen et al., 2002). Oxidative dissolution of sulfide minerals will release dilute sulfuric acid, iron, and other metal(loids). Subsequent neutralization of sulfuric acid by carbonate minerals, used in ABA to balance acid generation, results in the release of calcium, magnesium, and bicarbonate. Although silicate mineral weathering rates are slower than those of carbonates, silicates

predominate in coal overburden and can be major sources of calcium, magnesium, sodium, potassium, aluminum, and silicon (Brady et al., 1998; Clark et al., 2018; Hammarstrom et al., 2009). Dissolution of salts or *in situ* brines retained in the rock also releases sodium, calcium, sulfate, and chloride (as well as, barium, strontium, and bromide).

Laboratory-scale column and mesocosm-scale lysimeter experiments have been used to predict TDS release from coal industry materials (overburden, refuse, combustion byproducts) for nearly three decades (e.g., Brady et al., 1998; Clark et al., 2018; Daniels et al., 2014a; Daniels et al., 2016; Daniels et al., 2014b; Hornberger et al., 2004; Orndorff et al., 2015). These studies have shown that (1) rock type and extent of weathering influence TDS release, (2) shales and mudstones release more TDS than sandstones, and (3) unweathered rocks release more TDS than weathered materials. The specific conductance (SC) of column leachates typically starts near peak values, decreases during the first few pore volume leach cycles, and then stabilizes over the remaining leach cycles. For weathered overburden materials, peak leachate SC was often less than 500 µS/cm, a proposed regulatory limit (Cormier et al., 2013a, b). Coal refuse produced during coal processing generated considerably higher peak SC and associated concentrations of TDS, acidity, and major and trace elements compared to overburden and interburden strata removed during mining operations (Cravotta and Brady, 2015; Daniels et al., 2014b; Orndorff et al., 2015). In one study (Daniels et al., 2014b), columns (0.0012 m³ rock) were upscaled to mesocosms (1.5 m³ rock) using the same rock samples but with larger size fragments than in the columns. SC declined significantly in both the columns and the mesocosms. Compared to the columns, the peak leachate SC was higher and the temporal decline of SC was not as steady in the mesocosms likely because the mesocosms were operated in a less controlled environment. In either case, the use of such laboratory and field kinetic tests can involve months to years to obtain results and generally requires kilograms of rock materials.

A rapid batch extraction method that can be used to test small quantities of representative materials and that correlates well with field leachates would be of value to identify and manage rock types that release high TDS, and to characterize TDS from different energy extraction activities. For in-field determinations, "rapid" could refer to days if samples were shipped to a commercial laboratory, or hours if rock samples could be crushed and sieved, reacted with extractant solutions, and quantified for TDS release based on SC using a portable meter (discussed below). Because ABA parameters such as maximum potential acidity (MPA) and net neutralization potential (NNP) are used by coal companies to characterize overburden and are readily measured, their ability to be used to predict TDS release has been evaluated. Odenheimer et al. (2015) demonstrated that MPA and NNP may be useful to indicate general levels of low, moderate, and high TDS release; however, their semi-quantitative model was based on TDS computed from paste SC for a pulverized rock sample and did not consider upscaled or field-measured leachate characteristics. Modifying a method described by Barnhisel and Harrison (1976) and O'Shay et al. (1990), Orndorff et al. (2010) developed an alternative to the MPA method that used hydrogen peroxide (30% H_2O_2) to oxidize sulfide minerals. They found that the peroxide potential acidity (PPA) was better than MPA as a predictor of TDS release from low-S rocks. However, the potential application of PPA to predict TDS release from a wide range of rock types was not evaluated.

The objectives of this research were to 1) develop and test a batch extraction method to predict TDS release from a range of rock types associated with energy extraction, 2) compare different batch extraction methods to results reported for column leaching tests and field-scale leachate, 3) evaluate those batch extraction methods to determine the most reliable method to quantify TDS release, and 4) identify tracers in leachate that may distinguish coal mining-derived TDS from other energy extraction sources.

2. Materials and methods

2.1. Sample collection and preparation

A total of 65 sedimentary rock and coal samples were obtained from multiple sources (Table 1). The samples represent important fossil fuel-bearing strata in the Appalachian Basin, including bituminous coals and gas-producing shales. Eleven rock samples (3 weathered coal refuse, 8 overburden) were provided by Virginia Tech. Equivalent samples had been previously used in laboratory-scale, flow-through unsaturated column leaching experiments (Daniels et al., 2014b, 2016; Orndorff et al., 2015). Six rock samples (5 overburden, 1 pyrite) were provided by the U.S. Geological Survey.

(USGS). The five overburden samples had been previously characterized and used in laboratory-scale, flow-through column leaching experiments (Hammarstrom et al., 2009; Hornberger and Brady, 2009). The pyrite sample collected from the Bald Eagle Formation during construction of I-99 at the Skytop roadcut in Centre County, PA, along with paired water chemistry measurements had been previously described (Hammarstrom et al., 2005). Four overburden samples were provided by the University of Kentucky. Three of these sample materials had been used in field-scale lysimeter studies (mesocosms) (Agouridis et al., 2012; Sena et al., 2014) and all four had also been used in laboratory-scale, flow-through column leaching experiments (Daniels et al., 2016). Twenty unleached coal refuse samples from the roof and floor of the Pittsburgh coal, which is described by Brady et al. (1998), were collected from drill core materials stored in a repository maintained by the Pennsylvania Bureau of Topographic and Geologic Survey (PA TGS; Harrisburg, PA). Samples were collected from cores 8009, 8011, 8012, and 8013 that were drilled in Greene County, PA. One unleached Marcellus Shale sample was collected from drill core materials (Sullivan core at 8276 feet) stored by PA TGS. One unleached Utica/Point Pleasant Shale sample was collected from drill core

materials stored in a repository maintained by the Ohio Geologic Survey (Columbus, OH). Eight samples of Utica/Point Pleasant Shale drill cuttings were provided by two gas development companies working in Pennsylvania. Finally, in March 2017, a total of 11 weathered coal refuse samples, 3 coal samples, and 4 coal refuse leachate samples were collected from two coal refuse disposal facilities (referred to as Mine A and Mine B) in western Pennsylvania.

2.2. Rock type categorization

Rock samples were sorted into six operational categories: coal (n = 3), overburden (n = 17), weathered coal refuse (n = 14), unleached but oxidized coal refuse (n = 20), gas-bearing shale (n = 10), and pyrite (n = 1) (Table 1). Coal refuse and overburden categories were differentiated based on the definitions in Pennsylvania Code Title 25 (Environmental protection), Chapter 87 (Surface mining coal), Section 87.1 (Definitions) (25 Pa. Code § 87.1) (Commonwealth of Pennsylvania, 2018a). Specifically, overburden is defined as "the strata or material overlying a coal deposit or between coal deposits in its natural state and shall mean material before or after its removal by surface mining". Coal refuse is defined as "any waste coal, rock, shale, slurry, culm, gob, boney, slate, clay and related materials, associated with or near a coal seam, which are either brought aboveground or otherwise removed from a coal mine in the process of mining coal or which are separated from coal during the cleaning or preparation operations". Shales closer in age and stratigraphic position to coal formations were included in coal refuse or overburden categories. The gas-bearing shale category included only the Utica/Point Pleasant Shale or Marcellus Shale samples. Pyrite included one sample from the Bald Eagle Formation at Skytop roadcut (Hammarstrom et al., 2005).

2.3. Operational extractions

Once received, rock samples were freeze-dried using a Labconco FreeZone 4.5 freeze dry system until constant weight was attained. Samples were crushed to <4.75-mm using a hydraulic press at 44.5 kN and thereafter with a mortar and pestle until all particles were < 2-mm in diameter. Samples were then pulverized using a Spex 8000 ball mill to produce particles <0.5-mm diameter (passed through No. 35 sieve).

Pulverized rock samples were sent to Geochemical Testing, a certified commercial laboratory in Somerset, PA, to conduct three operational extractions and analyze the SC, pH, and solute concentrations of various leachates. A fourth extraction was conducted at Pennsylvania State University to measure strontium isotopes ($^{87}Sr/^{86}Sr$). In Leach 1 (L1), rock samples were reacted with distilled deionized water (DI) under an ambient atmosphere. In Leach 2 (L2), rock samples were reacted with DI water under a 10/90% CO_2/N_2 atmosphere. In Leach 3 (L3), rock samples were reacted with 30% H_2O_2 (70% DI) under a 10/90% CO_2/N_2 atmosphere. Aside from the differences noted above, the operational procedure for generating leachates followed the same steps. First, 10.00 ± 0.05 g of pulverized rock (<0.5-mm sieve size) was added to an Erlenmeyer 125-mL flask followed by 20 mL of the extractant solution. Addition of DI water in L1 and L2 was done rapidly in one aliquot. Addition of H_2O_2 in L3 was done slowly by adding 1 mL at a time to reduce bubbling caused by oxidation reactions. Flasks were then placed on a shaker table inside a controlled atmosphere apparatus. The lid of the controlled atmosphere apparatus was left open for L1 or sealed for L2 and L3. For the sealed conditions, 10/90% CO_2/N_2 gas was constantly flushed through the apparatus. All extractions were shaken for 4 h at 50 rpm and 20–22 °C. After 4 h, each sample was filtered through a 0.45-μm cellulose acetate filter and pH and specific conductance (SC) of the filtrate were measured immediately (Oakton multiparameter PCTestr 35, calibrated with standards and buffers at 20–22 °C). The filtrate was transferred to a 100-mL volumetric flask and DI water was added to dilute the leachate to a final volume of 100 mL for analysis of elemental concentrations. Blank samples were prepared with

DI water or H_2O_2 and followed all steps described above.

2.4. Analytical methods

A suite of analytes were measured for each of the three leachates (Supporting Information Tables SI–1 – SI-3). SC and pH were measured with electrodes submerged in the undiluted leachate. Major elements (Al, Ca, Fe, K, Mg, Mn, Na, P, S, Si) were measured after dilution to 100 mL on a Thermo Scientific iCAP 7400 inductively coupled plasma optical emission spectrometer (ICP-OES). Minor elements (As, Ba, Co, Cu, Li, Mo, Ni, Pb, Se, Sr, Th, Tl, Ti, U, V, Y, Zn, Zr) and rare-earth elements were measured on an Agilent 7900 inductively coupled plasma mass spectrometer (ICP-MS). Anions (Br, Cl, F, NO_3, SO_4) were measured on a Dionex DX-120 ion chromatograph (IC). Total inorganic carbon (TIC) was determined by infrared detection after persulfate oxidation (ASTM, 2017). ABA parameters were determined by standard methods (Sobek et al., 1978). Total sulfur (S) was analyzed by dry combustion, and maximum potential acidity (MPA) was calculated by multiplying total S (%) by 31.25 to obtain g kg^{-1} $CaCO_3$ equivalent. Neutralization potential (NP) was determined by reacting samples with HCl and titrating the effluent with NaOH using methods of Noll et al. (1988), without modification to account for siderite (Skousen et al., 1997). Net neutralization potential (NNP) was calculated by subtracting MPA from NP; negative NNP values imply a potentially acid-producing sample. The above analyses were conducted at Geochemical Testing, Somerset, PA.

Mineralogy of the rock samples was characterized by X-ray diffraction (XRD) using a PANalytical X'Pert 165 PRO MPD X-ray diffractometer equipped with a PIXcel detector operated in a 1D scanning mode with all channels active. Samples were subjected to Cu K-α radiation from 5 to 70° (2θ) at 45 kV and 40 mA. Semi-quantitative analyses were performed using whole pattern fitting in Jade 2010 software from Materials Data Incorporated of Livermore, CA, in conjunction with reference files from the International Centre for Diffraction Data PDF4 database. Mineral detection limits were about 3% (m/m) and uncertainty in mineral fractions was ±5%.

2.5. Speciation modeling methods

The PHREEQC 3.0 aqueous speciation model (Parkhurst and Appelo, 2013) was used with input values for leachate data, corrected for dilution (100/volume leachate recovered), to estimate SC by methods reported by McCleskey et al. (2012) and Appelo et al. (2010) as described by Cravotta and Brady (2015). Input data to PHREEQC included the sample temperature, pH, and the mass concentrations (mg/L) of TIC, SO_4, Cl, F, Br, NO_3–N, P, Si, Ca, Mg, Na, K, Li, Fe, Mn, Al, Ba, Sr, and Zn in the filtered leachates. Both methods calculated SC using the same speciated cations and anions (H^+, Li^+, Na^+, K^+, Cs^+, NH_4^+, Mg^{2+}, Ca^{2+}, Sr^{2+}, Ba^{2+}, F^-, Cl^-, Br^-, SO_4^{2-}, HCO_3^-, CO_3^{2-}, NO_3^-, and OH^-), trace metals (Al^{3+}, Fe^{2+}, Fe^{3+}, Mn^{2+}, and Zn^{2+}), and charged ion pairs (HSO_4^-, $NaSO_4^-$, $NaCO_3^-$, and KSO_4^-), however, the computations used to determine ionic conductivities were different. Briefly, the Appelo et al. (2010) method calculates the ionic conductivity of solute species using ion diffusion coefficients while the McCleskey at al. (2012) method calculates ionic molal conductivities using transport numbers. Both methods sum the ionic conductivity contributions to indicate the solution SC. Additional details on the SC computations are provided in the Supporting Information.

The concentration of total dissolved solids (TDS) was computed as the sum of the input concentrations of major dissolved constituents (Ca, Mg, Na, K, SO_4, Cl, CO_3, NO_3, SiO_2) (Fishman and Friedman, 1989) plus minor constituents (Sr, Ba, Fe, Al, Mn, Br), assuming that Fe, Al, and Mn formed hydrous oxides (FeOOH, AlOOH, MnOOH) instead of anhydrous compounds. Cravotta and Brady (2015) showed that TDS values computed accordingly were comparable to the laboratory-measured residue on evaporation at 180 °C for mine effluent samples. Osmotic pressure (OP) was computed as the sum of molal concentrations of the

same aqueous species used for SC calculations. The OP computation assumes that 1 mol/kg of each ion exerts approximately 1 mOsm/kg osmotic pressure (Haynes, 2013). Cravotta and Brady (2015) showed that computed values of OP for mine effluent samples were comparable to standard laboratory-measured values of OP using freezing point depression (Kiyosawa, 2003).

2.6. Data for upscaling comparisons

Several of the rock samples characterized by operational batch extractions were previously used in laboratory-scale flow-through column experiments or obtained from field sites with paired water samples (Table 1). For upscaling batch extractions to column experiments, 16 overburden and 3 weathered coal refuse samples were compared using mass-normalized TDS. As TDS was not reported for column experiments (only SC), a SC-to-TDS conversion factor (CV = TDS/SC = mg TDS L^{-1}/μS cm^{-1}) was calculated for each paired sample using the SC measured in L1 and the corresponding TDS value calculated using the input concentrations for PHREEQC (Supporting Information Tables SI–4). The cumulative TDS generated in the column experiment following approximately 14 or 40 discontinuous leaching events was calculated according to:

$$Cumulative\ column\ leached\ TDS \left(\frac{mg\ TDS}{L}\right) = \frac{\sum_{n=1}^{i} (SC_i \times V_i \times CV)}{V_{total}}$$

Eq 1

where, SC_i = SC measured from i-th leach event (μS cm^{-1}); V_i = volume of each leach event (L); CV = rock-specific SC-to-TDS conversion factor (mg TDS L^{-1}/μS cm^{-1}); and V_{total} = total volume of leaches (L).

Comparisons between field sites and batch experiments were made based on SC, as this parameter was reported for all field samples. A total of 10 field sites were included (referred to as Mine A, Mine B, Mines P, KY1, KY2, KY3, KY9, LKFC, BCS3, Skytop), where SC measured from a select number of rock samples were paired with a varied number of SC values measured in the field. For Mine A, 42 records of SC and additional analytes from leachate drains were compared to five weathered coal refuse samples collected from Mine A. For Mine B, 41 records from leachate drains were compared to six weathered coal refuse samples collected from Mine B. For Mines A and B, records were obtained from a field sampling event in March 2017 and from Hydrologic Monitoring Reports (HMRs) submitted by the coal companies to the Pennsylvania Department of Environmental Protection. For Mines P, three records from influent discharges to three Pittsburgh coal mining/processing plants on active underground mines (Cravotta and Brady, 2015) were compared to 17 unleached coal refuse samples stratigraphically adjacent to the Pittsburgh coal horizon (TGS1-TGS17, Table 1). Field results for KY1-KY3 are summarized by Sena et al. (2014). For KY1, 199 records from field lysimeters built on top of a valley fill were compared to rock sample KY1 (unweathered overburden). For KY2, 110 records from field lysimeters installed on top of a valley fill were compared to rock sample KY2 (unweathered overburden). For KY3, 203 records from field lysimeters installed on top of a valley fill were compared to rock sample KY3 (unweathered overburden). For KY9, 18,064 records from a leachate drain at the toe of two valley fills were compared to rock sample KY9 (unweathered overburden). Three records of untreated effluent from coal processing plants (Cravotta and Brady, 2015) and 24 records from discharges from abandoned mines in the lower Kittanning coal horizon (Cravotta, 2008) were compared to rock sample LKFC (unweathered overburden). Six records from influent discharges from active mines (Cravotta and Brady, 2015) and 10 records from discharge samples from abandoned mines in the lower to upper Freeport coal horizons (Cravotta, 2008) were compared to rock sample BCS3 (unweathered overburden). Four records from drainage from the Skytop roadcut collected in May 2004 (Hammarstrom et al., 2005) were compared to rock sample SKYPA (pyrite).
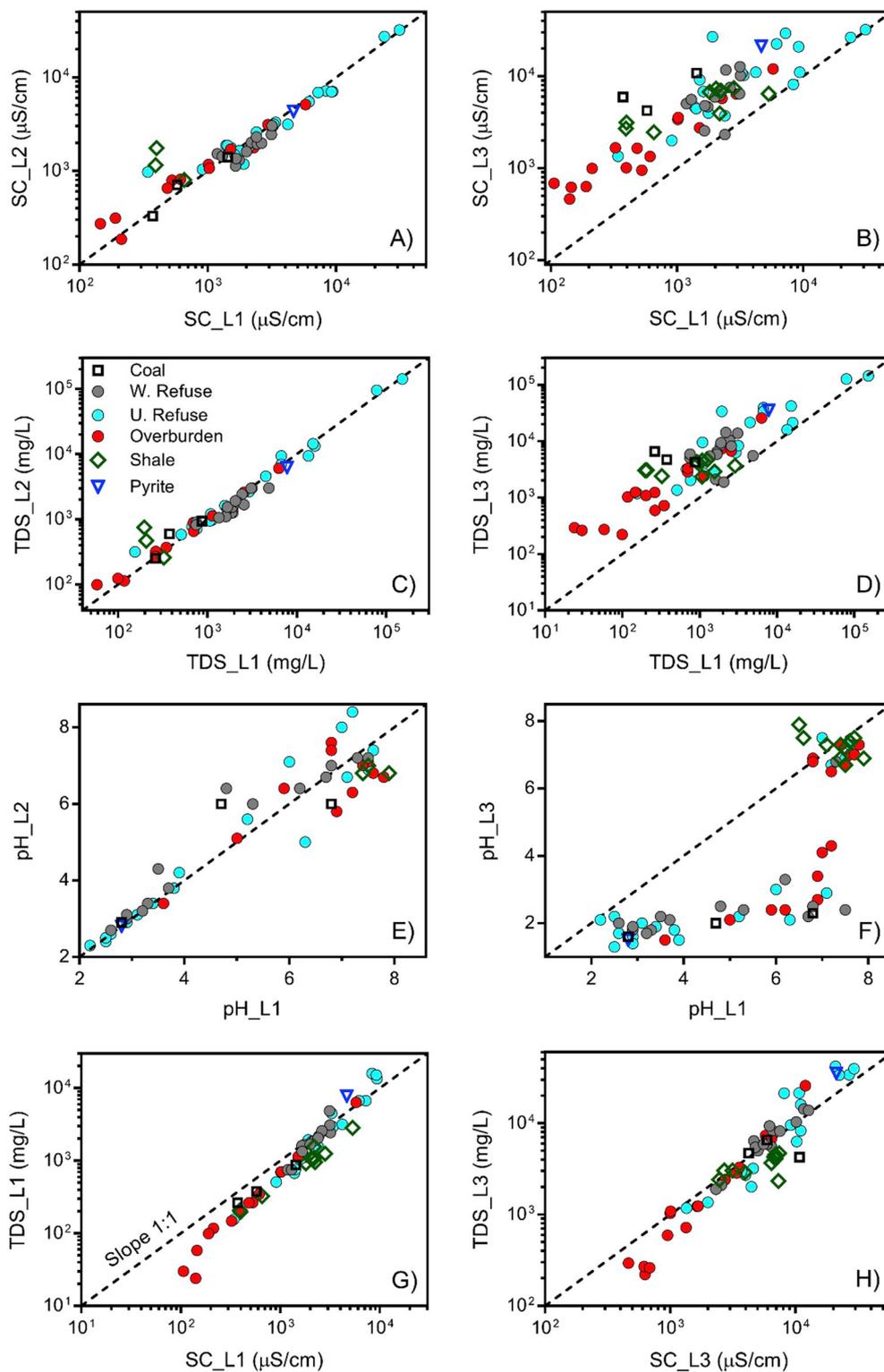
Fig. 1. Comparisons between different batch extractions. Panels A, C and E (n = 59) show correlations between Leach 1 (L1) and Leach 2 (L2). Panels B, D, and F (n = 63) show correlations between L1 and Leach 3 (L3). Panels G and H (n = 63) show correlations between SC and TDS in L1 and L3.

2.7. Statistical methods

Statistical differences between batch extractions (L1, L2, and L3) and upscaled results for selected parameters were evaluated using the Wilcoxon-signed rank test. Outliers were defined as values greater than the 75th percentile plus 1.5 times the inter-quartile distance, or values smaller than the 25th percentile minus 1.5 times the inter-quartile distance. SC and TDS were compared via correlation (Pearson) analyses. Comparisons between TDS from leaches were compared with TDS from column experiments using Pearson correlation and by comparing the fit of our data (R^2) with the line of equality y = x. SC data from the field were compared with SC from batch extractions using an unpaired *t*-test. Linear regression equations were generated for each rock category and for the full data set. R software was used for all statistical analyses.
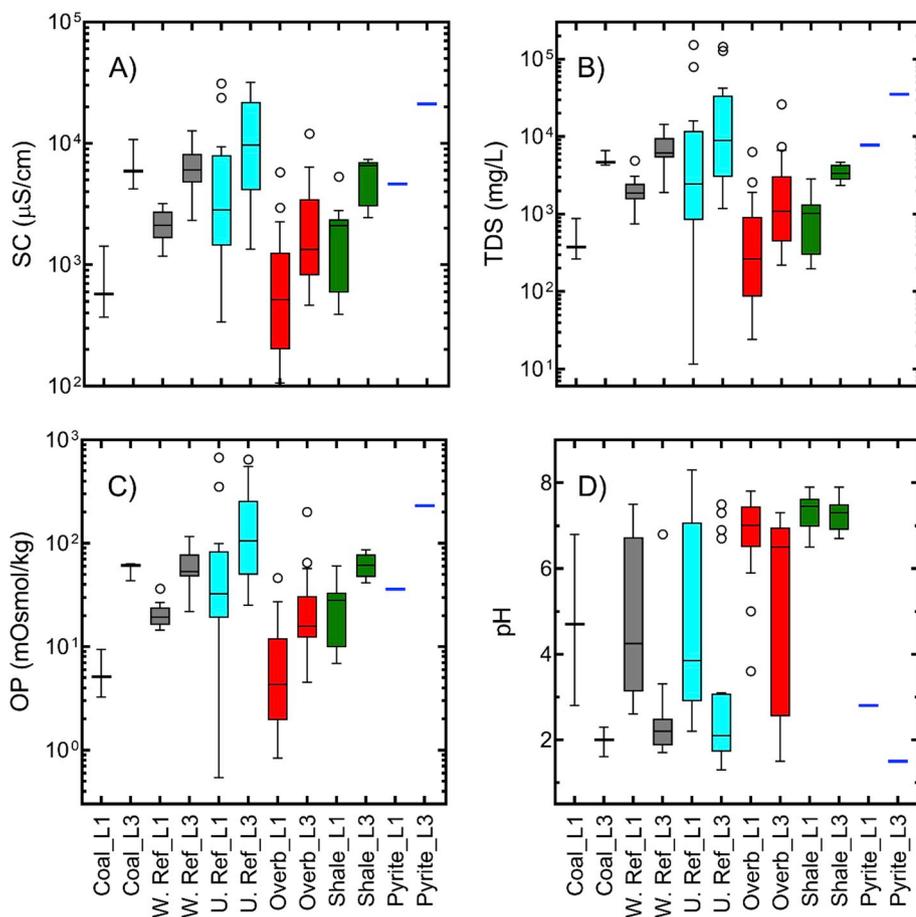
Fig. 2. Summary of chemistry for Leach 1 (L1) and Leach 3 (L3) for the six rock types. Coal (n = 3); Weathered Coal Refuse = W. Ref (n = 14); Unleached Coal Refuse = U. Ref (n = 20); Overburden = Overb (n = 17); Shale (n = 10); Pyrite (n = 1). Box plots show median, 25% and 75% quartile ranges. Whiskers show the minimum and maximum values. Outliers (circles) defined as any point at a distance greater than 1.5 times the interquartile range measured from the 75th to the 25th percentile.

3. Results and discussion

3.1. Comparison of operational extractions

Three operational extractions were designed to have varying reactivity with sulfides, carbonates, silicates, sulfates, and salts. DI water alone (L1) was presumed to extract weakly held exchangeable ions, salts, hydrolysis products, and high-solubility minerals. The equilibration of the DI extractant with 10% CO_2 atmosphere (L2) was hypothesized to promote carbonate dissolution. Although the pH of the L2 extractant blank was less than that of L1, as described below, this level of CO_2 did not create significant changes in the chemistries of leachates produced by L1 versus L2. In contrast, the 30% H_2O_2 in L3 promoted sulfide oxidation, and the production of sulfuric acid promoted the dissolution of many other minerals. Relationships between SC, TDS, pH, and TIC reflect the major reactions promoted by the extractant solutions. For example, for rocks with high sulfide and low carbonate contents (as determined by XRD), the increase in SC and TDS after the addition of H_2O_2 (i.e., L1 vs L3) was dominated by production of SO_4^{2-} and H^+. As another example, for rocks with low sulfide and high carbonate contents, the increase in SC and TDS after reaction with CO_2 (i.e., L1 vs L2) was controlled by release of Ca^{2+} and HCO_3^-.

Chemistry data for all rock samples are provided as spreadsheets in Supporting Information Tables SI–4 (leachates), SI-5 (blanks), and SI-6 (solids). For the blanks, the median (and range) of pH values were: L1, 6.0 (4.1–7.0, n = 7); L2, 5.1 (4.4–6.3, n = 5); and L3, 2.6 (1.8–4.5, n = 7). The median measured SC values for the blanks were 13 μS/cm (5.0–85, n = 7), 45 μS/cm (15–90, n = 5), and 227 μS/cm (33–973, n = 7) in L1, L2, and L3, respectively. The median calculated TDS values for the blanks were 26 mg/L (16–223, n = 7), 30 mg/L (14–34, n = 5), and 138 mg/L (26–339, n = 7) in L1, L2, and L3, respectively.

Operational extractions L1, L2, and L3 were compared based on pH, TDS and SC results from all rock samples (Fig. 1). The D'Agostino & Pearson normality test showed that the measured values for these parameters did not follow a normal distribution, therefore, comparisons between leachates were made using the Wilcoxon-signed rank test. This test showed that SC (measured and calculated), TDS, and pH values were significantly different (at 95% confidence) in L3 compared with L1 and L2, and that differences between L1 and L2 were not significant (Supporting Information Tables SI–7). In general, L3 showed an increase in SC and TDS and a decrease in pH compared with L1 (Fig. 1B,D,F). As noted above, this was the result of the oxidation of sulfide minerals promoted by the use of H_2O_2 in L3. The decreased pH promoted the dissolution of carbonate minerals and release of HCO_3^-, Ca^{2+}, and other ions into solution. Most of the samples showed a substantial increase in SC in L3 compared to L1 (Fig. 1B). However, six samples showed only a modest increase in SC (samples touching line of equality in Fig. 1B). These six samples contained high sulfate and low sulfide contents where the sulfate salts were quickly dissolved by water and the addition of H_2O_2 did not greatly enhance mineral dissolution. Of these six samples, four were unleached coal refuse (TGS 2A, TGS8, TGS10A, and TGD10B), one was weathered coal refuse (TNR2), and one was shale (SOH1). Consistent with statistical paired tests, results from L1 and L2 were similar (Fig. 1A,C,E). However, five samples showed an increase in SC in L2 compared to L1 (Fig. 1A). Of these five samples, two were gas-producing shales (SHM2, SHM3), one was unweathered coal refuse (TGS13), and two were sandstone overburden (VA16, WV5). The first three samples had abundant calcite and minor pyrite (Table 1). Although neither mineral was detected by XRD for VA16 or WV5 (Table 1), the two overburden samples had detectible NP and S (Table S3). In any case, the added CO_2 in L2 appears to have enhanced carbonate dissolution.
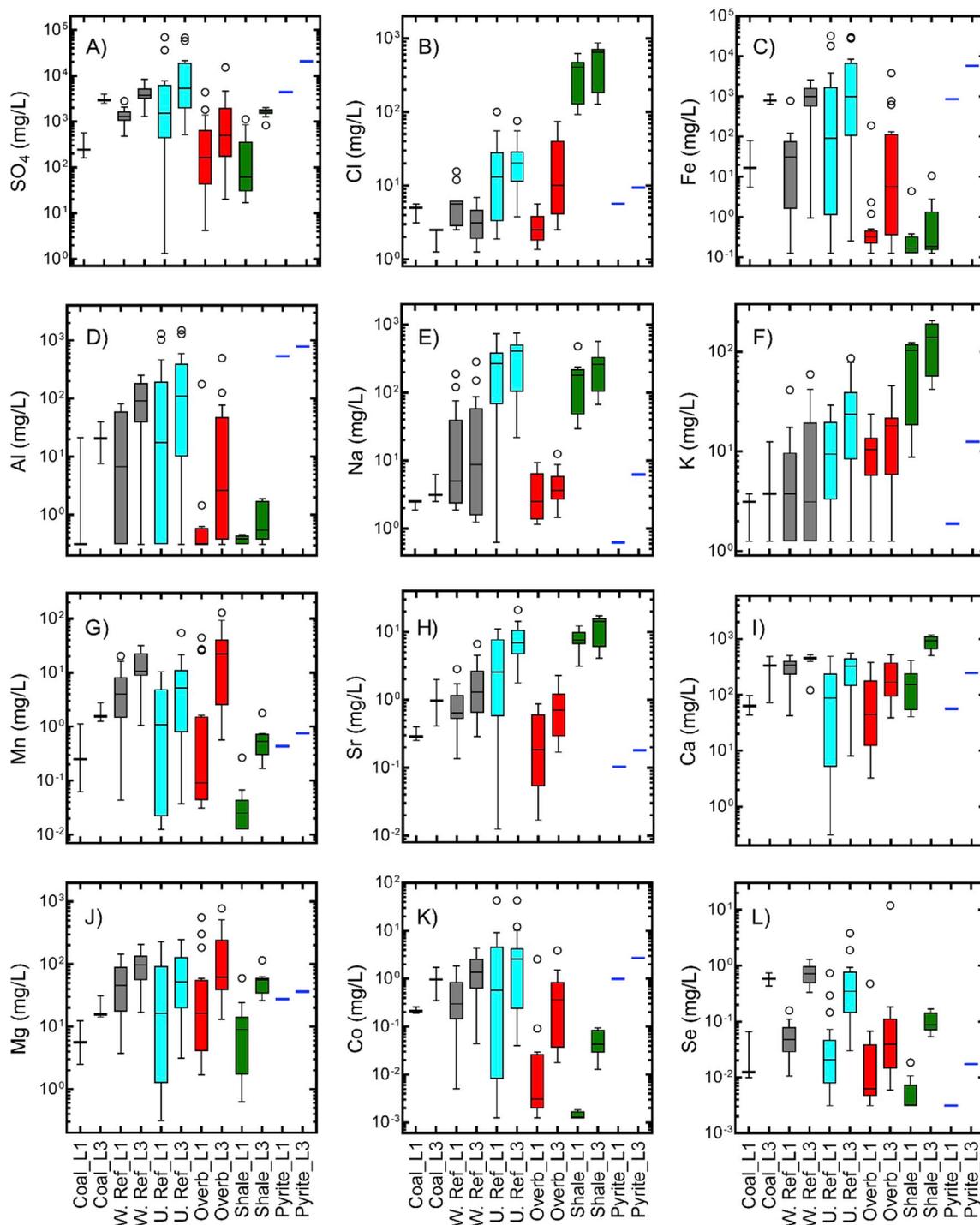
Fig. 3. Summary of chemistry for Leach 1 (L1) and Leach 3 for the six rock types. Coal (n = 3); Weathered Coal Refuse = W. Ref (n = 13 or 14); Unleached Coal Refuse = U. Ref (n = 20); Overburden = Overb (n = 17); Shale (n = 10); Pyrite (n = 1). Box plots show median, 25% and 75% quartile ranges. Whiskers show the minimum and maximum values. Outliers (circles) defined as any point at a distance greater than 1.5 times the interquartile range measured from the 75th to the 25th percentile. not been oxidized during core storage. Concentrations of transition metals and Se were similar between unleached and weathered coal refuse (Fig. 3).

Most pH values for L3 were lower than those for L1 (Fig. 1F) because of enhanced sulfide oxidation by $H_2O_2+CO_2$. However, many samples (23 of 65) with pH values ranging from 6.5 to 7.9 (samples clustered in upper right of Fig. 1F) exhibited little change in pH in L1 versus L3. These samples contained high carbonate and low sulfide contents, with corresponding positive values of NNP (Table 1 and Supporting Information Table S6) and produced enough alkalinity to neutralize the sulfuric acid produced. Of these 23 samples, 8 were shales, 6 were unleached coal refuse, 2 were weathered coal refuse, and 7 were

overburden.

For the six rock types, median values for pH, SC, TDS, and OP were used to compare L1 and L3 (Fig. 2). In general, pyrite and unleached coal refuse produced the highest median values for SC and TDS while overburden produced the lowest median values. Median values for these parameters from weathered coal refuse, coal, and shale always ranked in intermediate positions, although the order switched depending on the analyte or extraction method. For instance, the order for TDS (mg/L) in L1 was pyrite (7,770) > unleached coal refuse (2,430) > weathered coal
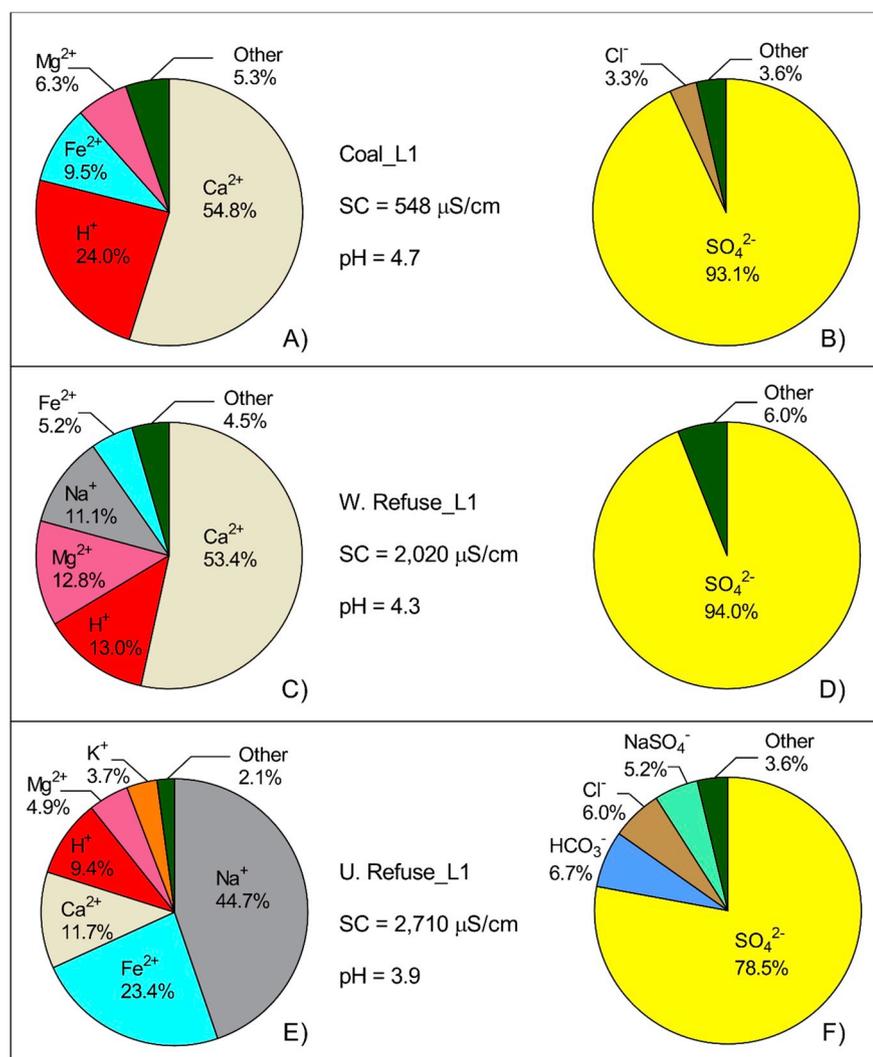
Fig. 4. Ionic contributions to specific conductance in Leach 1 (L1). Median specific conductance (SC), and ionic contributions calculated according to McCleskey et al. (2012). Median pH is shown for each rock category. Left panels show cationic contributions normalized to median SC for each rock category. Right panels show anionic contributions normalized to median SC for each rock category. Coal (n = 3); Weathered Coal Refuse = W. Refuse (n = 14); Unleached Coal Refuse = U. Refuse (n = 20); Overburden (n = 17); Shale (n = 10); Pyrite (n = 1).

refuse (1,870) > shale (1,020) > coal (375) > overburden (262), while the order in L3 was pyrite (35,200) > unleached coal refuse (8,920) > weathered coal refuse (6,160) > coal (4,700) > shale (3,360) > overburden (1,080). Median values of TDS and SC in L3 were all higher than corresponding proposed regulatory reference levels, 500 mg/L and 300–500 μS/cm (Cormier et al., 2013a,b; Pond et al., 2008; Timpano et al., 2010), respectively, for all rock types. Except for overburden, median values of OP in L3 were all higher than the regulatory reference level of 50 mOsmol/kg (Commonwealth of Pennsylvania, 2018b).

It is important to note that our results for TDS from Marcellus and Utica/Point Pleasant Shale are not consistent with values reported for waters produced from unconventional gas wells drilled into these formations. For instance, Tasker et al. (2019) reported median values of TDS of 240,000 mg/L for Utica/Point Pleasant Shale wells (n = 24) and 225,000 mg/L for Marcellus Shale wells (n = 60). Those values are nearly two orders of magnitude greater than the median TDS obtained for gas-bearing shales with our most aggressive extractant (TDS_L3 = 3360 mg/L). Produced water from hydraulically fractured shale gas wells could encounter extensive small fracture networks equivalent to exceptionally low water-to-rock ratios or could interact with brines that were not within (or preserved) in our samples. This is consistent with other studies that have shown TDS values from batch extractions of Marcellus Shale are much lower than corresponding field-produced waters (Chapman et al., 2012; Phan et al., 2015; Rowan et al., 2011;

Stewart et al., 2015; Tasker et al., 2016; Warner et al., 2012).

3.2. Leachate composition

The median concentrations of dissolved metals, metalloids, and anions varied based on rock type, weathering extent, and final pH of the extract (Fig. 3). Based on sample mineralogy and leachate composition, the principal mechanisms for TDS generation are oxidation of sulfide minerals (with production of H_2SO_4) that promotes increased solubility of metals (e.g., Fe, Al, Mn), dissolution and hydrolysis of carbonate and silicate minerals to neutralize acidity, and dissolution of high-solubility minerals such as sulfates and salts. Because of its high organic carbon content, coal contained relatively low concentrations of metals. Based on XRD (Table 1), coal samples contained quartz, calcite, clays, and pyrite. Concentrations of Si, Ca, Al, Fe, and SO_4 in the coal leachates are consistent with this mineral assemblage.

Weathered coal refuse had been exposed to shallow subsurface weathering for years, while rock cores classified in this study as unleached coal refuse had been exposed to humid air only while archived in core boxes. These differences in weathering extent led to distinct differences in leachate chemistry. Unleached coal refuse released higher alkali metals, notably Na, and higher Cl compared to weathered coal refuse (Fig. 3), reflecting that salts had been preserved in storage. Plots of Cl versus Na molar concentrations in both L1 and L3
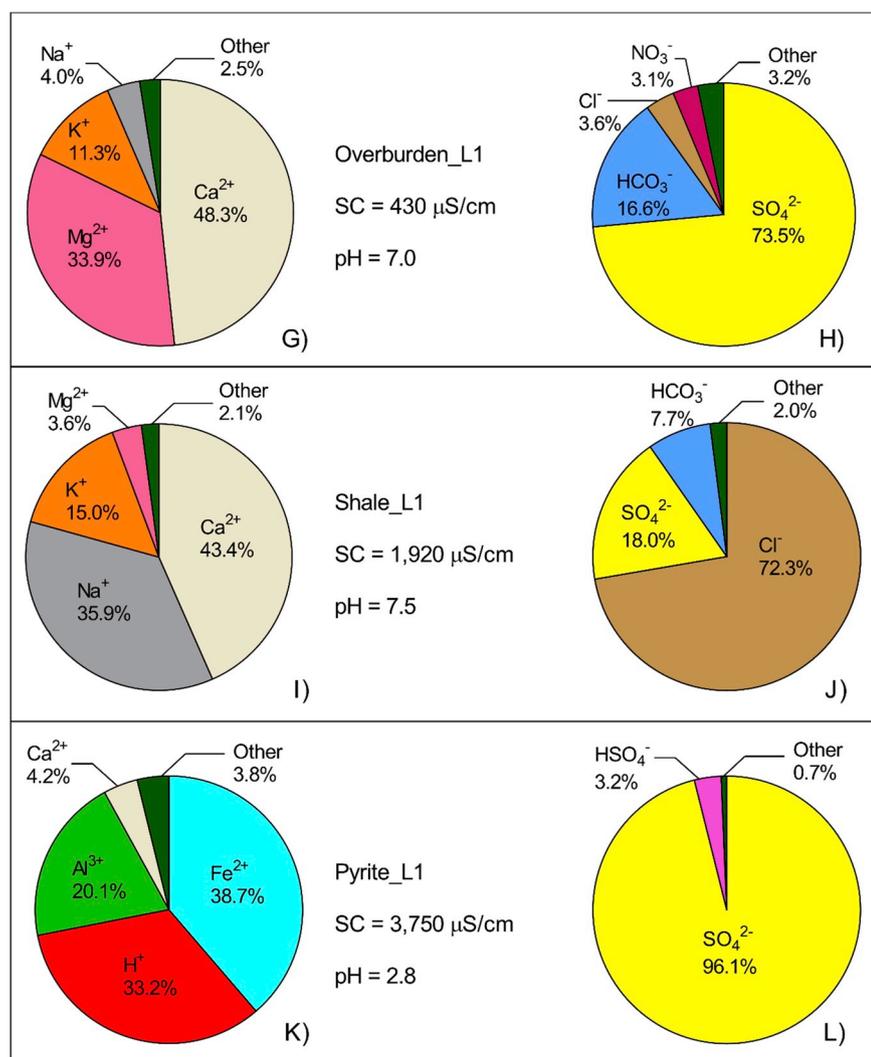
Fig. 4. (*continued*).

showed that only the shale samples consistently plotted along the 1:1 line of equality supporting the assumption of NaCl dissolution (Supporting Information Fig. SI-1). Na was also likely sourced from exchange reactions and silicate neutralization. Unleached coal refuse also released higher SO_4, notably in L3, compared to weathered coal refuse, reflecting that some sulfides had not been oxidized during core storage. Concentrations of transition metals and Se were similar between unleached and weathered coal refuse (Fig. 3).

Differences in mineral composition (Table 1) of overburden, shale, and pyrite help explain differences in leachate chemistry. In L3, overburden samples released low alkali metals and chloride, reflecting low entrained salt content, and low SO_4 reflecting low sulfide content. Gas-bearing shale samples released high alkali metals and the highest amounts of Cl and Br, reflecting relatively high salt content, low SO_4 reflecting low sulfide content, and high alkaline earth metals, notably Sr and Ba, reflecting high carbonate content. The sole pyrite sample released the highest amounts of Fe, SO_4, Al, Pb, Zn, and As reflecting high sulfide content.

3.3. Contribution of ionic species to specific conductance

The relative contributions of ionic species to the SC were calculated for all rock types in L1 and L3 using the method of McCleskey et al. (2012). For L1, the major cationic contributions to SC were $Ca^{2+} > Na^+ > Mg^{2+} > Fe^{2+} > H^+ > K^+$, and major anionic contributions to SC were $SO_4^{2-} > Cl^- > HCO_3^-$ (Figs. 4 and 5). However, the rank of ion contributions to SC differed slightly depending on rock type. For instance, Ca^{2+} was the dominant cation in coal, weathered coal refuse, overburden, and gas-bearing shale, but Na^+ was the dominant cation in unleached coal refuse and Fe^{2+} was the dominant cation in pyrite. On the other hand, the anionic contributions to SC were dominated by SO_4^{2-} for all rock types except for gas-bearing shale, where Cl^- was most abundant. The high contribution of Cl^-, Na^+, and Ca^{2+} to SC in gas-bearing shales (Fig. 4I and J) provide evidence for salt dissolution. The general contribution of principal cations and anions to SC in L1 were consistent with the mechanisms of TDS generation discussed above.

With the addition of H_2O_2 to L3, sulfide oxidation and the consequent release of sulfuric acid became an important mechanism for ion mobilization by mineral dissolution. While the major ions that contribute to SC in L3 were similar to those in L1, the contribution of H^+ increased markedly in L3 (Fig. 5). The major cationic contributions to SC in L3 were $H^+ > Ca^{2+} > Fe^{2+} > Na^+ > Mg^{2+} > K^+$. Anionic contributions to SC in L3 were dominated by $SO_4^{2-} > HSO_4^- > Cl^- > HCO_3$. Na^+ was an important contributor to SC from unleached coal refuse and gas-bearing shale. The increased release of Na^+ from gas-bearing shale with L3 compared to L1 is consistent with silicate mineral decomposition combined with salt dissolution. Cl^- was an important contributor to SC of L3 only from gas-bearing shale, where the SO_4^- release was greater than Cl^-. Mg^{2+} and HCO_3^- were important contributors to SC only from overburden. Although pH of L3 remained near-neutral for the gas-
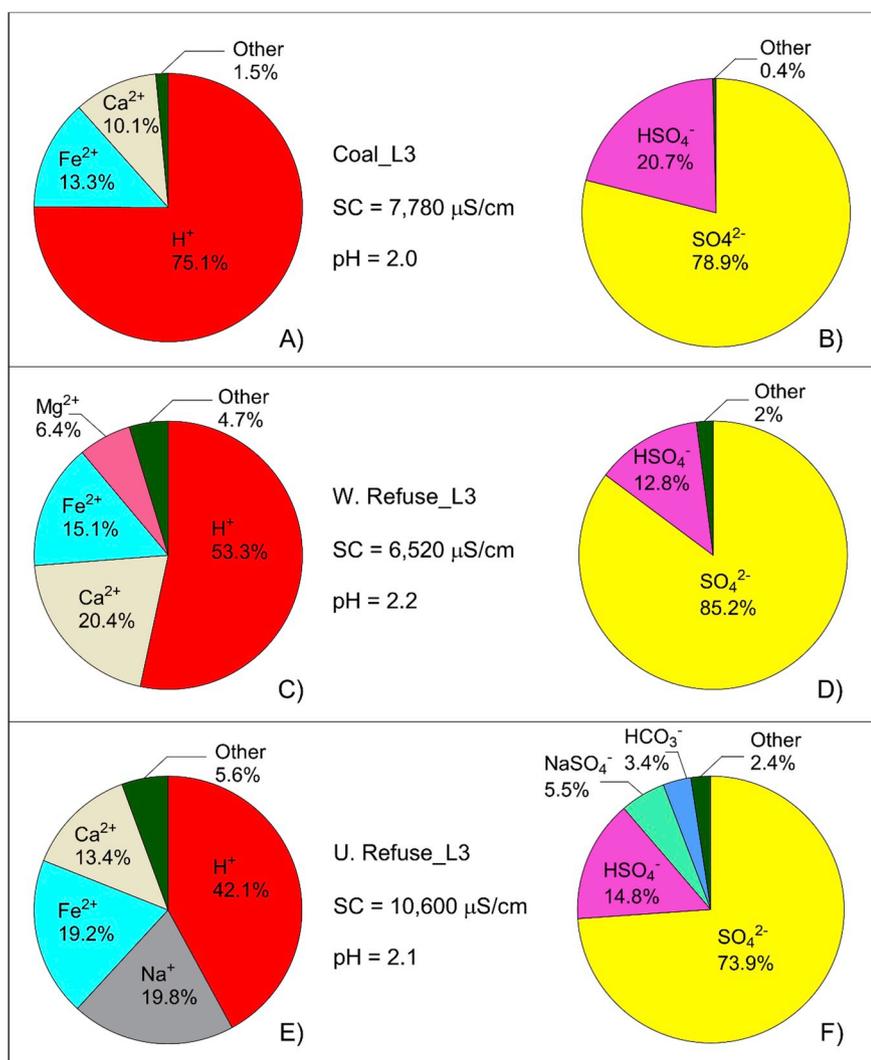
Fig. 5. Ionic contributions to specific conductance in Leach 3 (L3). Median specific conductance (SC), and ionic contributions calculated according to McCleskey et al. (2012). Median pH is shown for each rock category. Left panels show cationic contributions normalized to median SC for each rock category. Right panels show anionic contributions normalized to median SC for each rock category. Coal (n = 3); Weathered Coal Refuse = W. Refuse (n = 14); Unleached Coal Refuse = U. Refuse (n = 20); Overburden (n = 17); Shale (n = 10); Pyrite (n = 1).

bearing shale and overburden, the increased release of SO_4^{2-}, Ca^{2+}, and Na^+ with L3 compared to L1 demonstrates the importance of mineral decomposition in conjunction with pyrite oxidation, acidification, and neutralization.

3.4. Upscaling from lab to field

An important part of this study was to test the capability of the proposed rapid batch extractions on 10-g samples for predicting TDS release from coal refuse and overburden of larger size and at longer time scales. Available information from previous column studies and water quality data from 10 field sites were compared with our batch extractions. Seventeen overburden samples and three weathered coal refuse samples (Table 1) were previously analyzed in column studies (Agouridis et al., 2012; Daniels et al., 2009, 2014b, 2016; Hornberger and Brady, 2009; Odenheimer et al., 2015; Sena et al., 2014). In general, all column studies maintained unsaturated conditions with simulated rainfall (pH 4.6) leaching events. The number and frequency of leach events, the rainfall volume, and the rock mass in the columns varied between experiments conducted by Daniels et al. (2016) versus Hornberger and Brady (2009). Daniels et al. (2016) constructed columns with 1200 cm³ (~1800 g) of rock and applied 125 mL of synthetic rain water twice a week for a total of 40 leach events. Hornberger and Brady (2009) constructed columns with 1300 to 2100 g of rock and applied 190–650

mL of synthetic rain water once a week for a total of 14 leach events. As described below, the overall water volume (sum of individual leaches) to rock mass ratio strongly controls leachate chemistry. As these two groups of researchers ultimately used similar water-to-rock ratios in their column studies, results from all studies are comparable when normalized to rock mass. In general, SC in the column leachates started at the highest values, declined in an exponential manner, and then approached an asymptotic minimum. Comparisons with batch experiments were made based on cumulative TDS calculated from the column experiments. Column leachate data were compiled as SC and then converted to TDS, based on rock-specific correlations derived from L1, using Eq (1). Based on this approach, we found that TDS from batch extraction L1 and cumulative TDS calculated from column experiments were well correlated (Fig. 6). These batch and column methods likely produced similar amounts of cumulative TDS because the water-to-rock ratios used in all experiments were of similar order of magnitude (2 mL-to-1 g in batch experiments versus 2.5 mL-to-1 g to 3.2 mL-to-1 g in column experiments), and the smaller particle size used in the batch extractions (≤0.5-mm for batch experiments versus ≤ 2-mm to ≤ 1.25-cm for column experiments) may promote more rapid release of TDS. However, overburden samples that produced the lowest amounts of TDS in L1 produced more cumulative TDS via column leaching (lower left of Fig. 6A).

TDS from batch extraction L3 tended to overpredict cumulative TDS
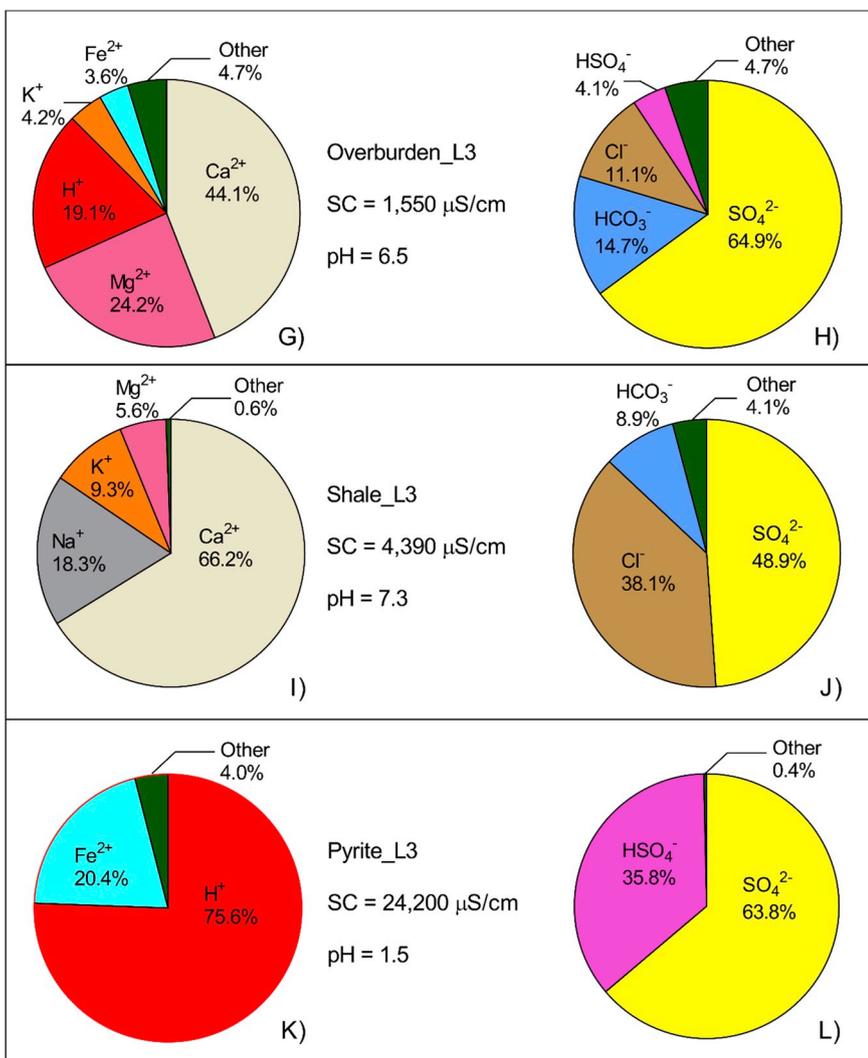
Fig. 5. (continued).

Overburden_L3
SC = 1,550 μS/cm
pH = 6.5

Shale_L3
SC = 4,390 μS/cm
pH = 7.3

Pyrite_L3
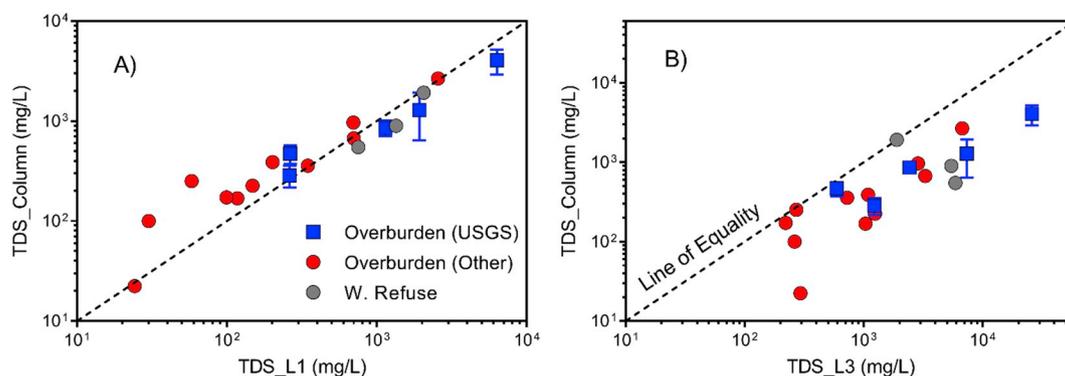SC = 24,200 μS/cm
pH = 1.5

Fig. 6. Cumulative total dissolved solids (TDS) calculated for column experiments (calculated using Eq. (1)) versus TDS measured in Leach 1 (L1) and Leach 3 (L3). For column experiments with replicates, symbols represent mean values and error bars represent standard deviation. Error bars smaller than the symbol size are not shown. Weathered coal refuse = W. Refuse (n = 3). Overburden samples donated by USGS = Overburden (USGS) (n = 5), overburden samples from other sources = Overburden (Other) (n = 11).

calculated from the column experiments (Fig. 6B; Supporting Information Tables SI–8). These results suggest that even multiple column leaching events cannot achieve the extractive strength of $H_2O_2+CO_2$ used in batch extraction L3. Furthermore, these results suggest that multiple discontinuous rainfall leaching events do not substantially increase the extractive strength of synthetic rain (or physical access to

additional reactive sites), and that the cumulative water-to-rock ratio exerts greater control on leachate chemistry for water extractions.

In contrast to the column experiments, field SC was better correlated to SC from batch extraction L3 versus L1 (Fig. 7; Supporting Information Tables SI–9). Field data and batch extractions were compared for 9 of the 10 field sites using an unpaired t-test (not enough data were
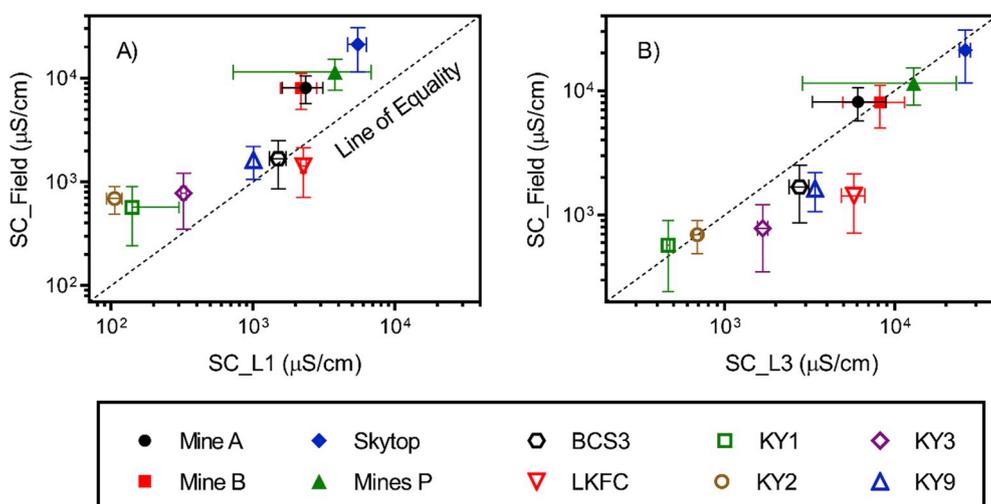
Fig. 7. Relationships between specific conductance (SC) measured in the field and in A) Leach 1 (L1) and B) Leach 3 (L3). Symbols represent mean values and error bars represent standard deviation. For SC measured in the field: Mine A (n = 42); Mine B (n = 41); Mines P (n = 3); KY1 (n = 199), KY2 (n = 110), KY3 (n = 206), KY9 (n = 18,064); LKFC (n = 25); BSC3 (n = 16); Skytop (n = 4). For SC measured in L1: Mine A (n = 5); Mine B (n = 6); Mines P (n = 17); KY1 (n = 3); KY2 (n = 3); KY3 (n = 3); KY9 (n = 1); LKFC (n = 3); BSC3 (n = 3); Skytop (n = 8). For SC measured in L3: Mine A (n = 5); Mine B (n = 6); Mines P (n = 17); KY1 (n = 3); KY2 (n = 3); KY3 (n = 3); KY9 (n = 1); LKFC (n = 2); BSC3 (n = 2); Skytop (n = 14).

Table 2

Correlations between TDS and ABA parameters. MPA = Maximum Potential Acidity; SC=Specific conductance; r = correlation coefficient.

Rock Type	N	Linear regression	Pearson Correlation	
		SC_L3 = m*MPA + b	R	P
Leach 3				
Coal	3	SC_L3 = −89.6*MPA +11700	−0.563	0.619
W. Refuse	14	SC_L3 = −7.74*MPA +7280	−0.187	0.523
U. Refuse	17	SC_L3 = 75.6*MPA +4010	0.835	<0.0001
Overburden	13	SC_L3 = 86.5*MPA +1220	0.921	<0.0001
Shale	10	SC_L3 = 74.0*MPA +2350	0.387	0.269
All data	58	SC_L3 = 57.9*MPA +3350	0.663	<0.0001
Rock Type	**N**	**Linear regression**	**Pearson Correlation**	
		TDS_L3 = m*MPA + b	**R**	**P**
Coal	3	TDS_L3 = 56.8*MPA +2170	0.989	0.0949
W. Refuse	14	TDS_L3 = −8.41*MPA +8090	−0.167	0.568
U. Refuse	17	TDS_L3 = 130*MPA − 2770	0.858	<0.0001
Overburden	13	TDS_L3 = 196*MPA − 267	0.985	<0.0001
Shale	10	TDS_L3 = 26.1*MPA +2420	0.326	0.358
All data	58	TDS_L3 = 100*MPA − 1880	0.692	<0.0001

available to test the other three sites). Seven of the nine sites showed no statistical differences between field SC and L3 SC. We note that field SC values from rock disposal facilities (that is, excess spoil fills pursuant to SMCRA) change over time, and that the 'age' of the rock/leachate could not be controlled in this study. However, we chose to analyze all these sites together because they represent the largest, most analytically consistent data set available for addressing our research objectives. In contrast, five of the nine sites showed significant statistical differences between field SC and L1 SC, where SC results from L1 underestimated the field SC.

Collectively these upscaling comparisons suggest that weathering in the field is influenced by acid-formation and neutralization reactions that produce greater solute concentrations than simple dissolution of soluble salts and exchangeable ions by water alone. Column leaching experiments produce high SC in the first leaches, but values decline rather quickly. Scaling up from column experiments to field sites is challenging for a number of reasons. One obvious issue is that coal refuse disposal fills often contain millions of cubic meters of rock such that the rock-to-water ratio is dramatically greater in the field as compared to tens of pore volumes eluted through laboratory columns. Unlike column experiments, water percolating through rocks stored in disposal fills may encounter multiple and much longer flow paths such that the water encounters more 'fresh' reactive material. Water may migrate through these rocks much slower and encounter many more wetting-and-drying

cycles as compared to column experiments such that the field leachates oxidize more sulfides, generate a lower pH, and solubilize more metals. Rocks in disposal fills may also disaggregate over long periods of time, effectively increasing the rock-to-water ratio.

3.5. Correlations between SC and ABA parameters

Acid-base accounting (ABA) parameters are used to identify and segregate rocks with high acid generation potential (or blend with rocks with high alkalinity). The use of ABA parameters to identify TDS release potential would be cost-effective for coal mine operators, provided that the TDS predictions based on ABA were accurate. Of all the correlations between ABA parameters and analytes measured in L1 and L3 extracts, maximum potential acidity (MPA) displayed the most promising correlations with SC from L3 (Table 2). It was anticipated that MPA + NP might better predict TDS release, but this did not produce an improved relationship. While MPA is certainly correlated with TDS release, correlation coefficients were not high for all rock types and notably low for weathered coal refuse, plus different linear regression coefficients (slope and intercept) were indicated for different rock types (Supporting Information Fig. SI-2). Therefore, an additional measure such as batch extraction L3 to measure TDS release potential would still be needed.

Correlations between XRD-based mineral contents, ABA parameters, and selected leachate chemistry parameters (Supporting Information Tables SI–11) confirmed that samples containing sulfide and sulfate minerals had higher total S, and samples containing calcite and dolomite had higher NP. The strongest predictor of leachate salinity (SC, TDS, or OP) was the total S content and the presence of sulfur minerals. These correlations support the hypothesis that L1 liberates sulfur and iron mainly from iron sulfate minerals formed by prior oxidation of pyrite. Identification of specific iron sulfide or sulfate minerals does not seem to be particularly informative for predicting water chemistry; however, oxidation of the sulfides is required to release the potential acidity. Consequently, the significance of correlations between ABA parameters and salinity parameters increases from L1 to L3 (which seems to mobilize greater Ca from carbonates in conjunction with sulfide oxidation).

3.6. Correlations among SC, pH, and element concentrations

Concentrations of major and minor elements and various measures of ionic strength were positively correlated with the measured SC (Fig. SI-3) and inversely correlated with the pH (Fig. SI-4). The apparent SC and pH distributions and correlations reflect the weathering contributions of different rock types (Fig. SI-5), with increased SC and decreased pH resulting from more extensive weathering of rock samples from coal-
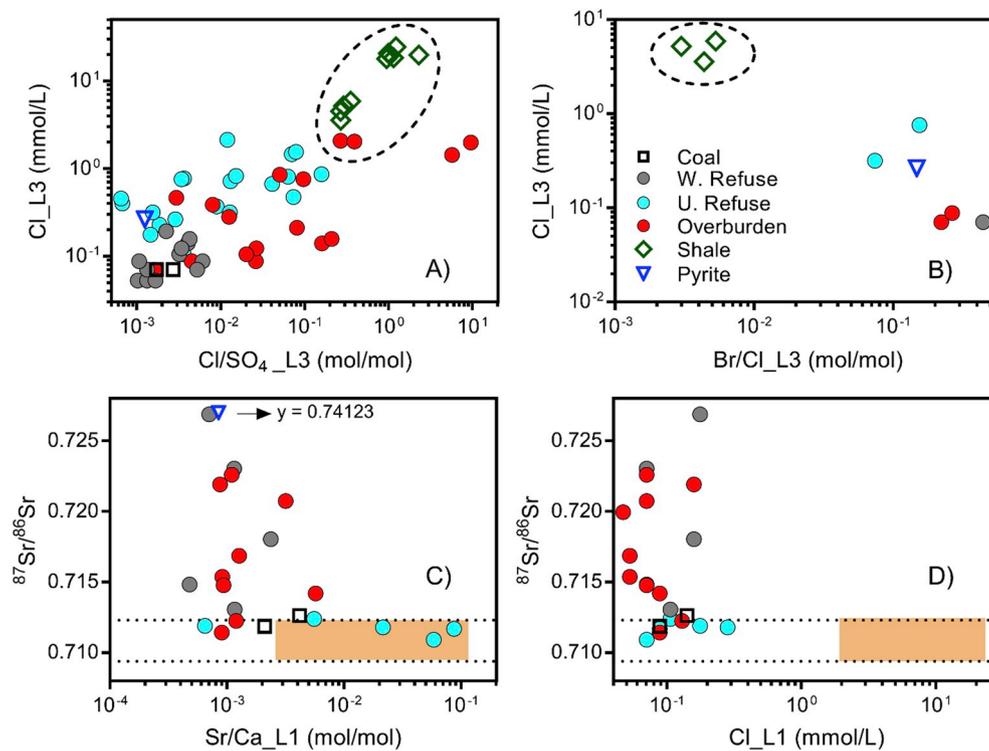
Fig. 8. Relations between sulfate or bromide and chloride to distinguish leachates of gas-bearing black shales from coal-associated rocks.

bearing strata containing pyrite and associated oxidation products; such samples produced the highest SC and the lowest pH. For Figs. SI-3 and SI-4, the DI and $H_2O_2+CO_2$ leaches (L1 and L2, combined) were binned based on SC interval (100–1000; 1000–2000; 2000–5000; 5000–10,000; 10,000–20,000; and 20,000–35,000 μS/cm) or pH interval (0.5–1.5; 1.5–2.5; 2.5–3.5; 3.5–4.5; 4.5–5.5; 5.5–6.5; 6.5–7.5; and 7.5–8.5) irrespective of rock type. The pH-dependency of concentrations of metals (e. g., Fe, Al, Cd, Co, Cu, Ni, Pb, Zn) reflected corresponding mineral or metal-oxide solubilities. The oxyanion-forming metalloids As and Se were also more soluble at lower pH and higher SC. Although the concentrations of all major and trace constituents generally increased with decreased pH, the concentrations of chloride, sodium, potassium, strontium, and barium also exhibited an increase from near-neutral pH to alkaline pH values. The increased concentrations of these elements with pH result from dissolution of constituents, including salts, in shale samples from gas-bearing strata (Figure SI-5). Thus, despite having limited acid-forming potential, the gas-bearing shales evaluated in this study can be sources of elevated salinity.

The ability to distinguish the source(s) of salinity becomes challenging in the Appalachian Basin where coal mining, conventional oil and gas (O&G) production, and unconventional gas production coexist. As noted above, shales produced Na–Cl waters that were generally distinct from Ca–SO$_4$ waters produced from coal-associated rocks. Additionally, bromide, strontium isotopes ($^{87}Sr/^{86}Sr$), and radium isotopes ($^{228}Ra/^{226}Ra$) have all been used to identify the addition of O&G-produced water into freshwater systems (Chapman et al., 2012; Rowan et al., 2011; Johnson et al., 2015; Warner et al., 2012). The leachate chemistry from gas-producing shales (nine samples from Utica/Point Pleasant Shale plus one sample from Marcellus Shale) compared to coal mining-associated materials evaluated in this study showed that differentiation with Cl versus Cl/SO$_4$ molar ratio could be effective for salinity source identification (Fig. 8). Br was not a robust tracer in this study because it was below detection in most samples from coal-bearing strata (e.g., 24 of 65 samples had measurable Br in L1; 9 of 65 samples had measurable Br in L3). Where Br values were above detection limits, Cl/Br and Cl were greater in the leachate from gas-bearing shale samples

compared to the coal-associated rocks.

4. Conclusions

Increased salinization of fresh water resources is a growing concern even in water-rich regions such as the Appalachian Basin. Management of activities and industries that release TDS could reduce salinity, SC, and other measures of ionic strength of receiving waters. For coal mining, segregation and isolation of rocks that produce high levels of TDS is one obvious management strategy. To implement this strategy, a rapid and simple method to identify these rocks by quantifying TDS release is required. In regions with coal mining and other sources of TDS (e.g., coal-bed methane, oil & gas development, road brining), source identification could also help reduce TDS release and enhance the information available to decision makers.

In this study, a rapid batch extraction method was evaluated to estimate potential for TDS release by 65 samples of rock from coal and gas-bearing strata of the Appalachian Basin in the eastern U.S. Three different extractant solutions were reacted with crushed rock samples (20 mL:10 g) for a duration of 4 h: deionized water (DI), DI equilibrated with 10% CO_2 atmosphere (DI + CO_2), or 30% H_2O_2 under 10% CO_2 ($H_2O_2+CO_2$). The TDS, SC, osmotic pressure, and ionic strength of the extracts were positively correlated and could be interpreted interchangeably. The pH of extractant blanks decreased in the order DI (6.0), DI + CO_2 (5.1), and $H_2O_2+CO_2$ (2.6), which indicated the 30% H_2O_2 was mildly acidic as well as an oxidant. The DI extractant was effective for mobilizing soluble SO_4 and Cl salts, which are predominant sources of TDS upon initial wetting of crushed rock. The DI + CO_2 extractant increased the weathering of carbonates present in some samples, but did not significantly increase TDS production compared to the DI extraction when considering the whole set of samples. The $H_2O_2+CO_2$ extractant increased the weathering of sulfides (and carbonates) and resulted in the greatest TDS production and lowest pH values.

Of the 65 samples, 19 had leachate chemistry data from flow-through column experiments and 35 were paired to 10 field sites with leachate chemistry data from multiple field records. When accounting for the

mass of rock-to-volume of extractant, TDS measured in batch extractions was strongly correlated to cumulative TDS calculated from column experiments. TDS measured in batch extractions using 30% H_2O_2 under 10% CO_2 was higher and poorly correlated to cumulative TDS calculated from upscaled column experiments. Results suggest that the cumulative water-to-rock ratio controls leachate chemistry in batch extractions using DI water or flow-through configurations using synthetic rain.

Because TDS was not reported or complete analyses of ions required to compute TDS were not available for numerous field samples, batch extractions and field measurements were compared based on SC. In contrast to column experiments, field SC was better correlated to SC measured from $H_2O_2+CO_2$ extractions versus DI extractions. The field SC and SC from $H_2O_2+CO_2$ extractions were statistically indistinguishable for 7 of 9 paired data sets while SC from DI extractions underestimated field SC in 5 of 9 cases. Compared to column leaching over months or waiting until mined rock begins weathering in the field, the rapid batch extractions of small samples are efficient and informative. The small sample size used in batch tests permits testing of specific lithologies or strata. Results were comparable among the rapid batch tests and longer-term laboratory or field data sets. Upscaling comparisons suggests that (1) weathering reactions in the field are more aggressive than DI water or synthetic rainwater extractants used in batch or column tests, and (2) a batch extraction method utilizing 30% H_2O_2 (which is mildly acidic without CO_2 enrichment) with corresponding measurements of SC could be effective for identifying rocks that will release high amounts of TDS. Additional measurements of pH, SO_4, and Cl in the extracts may be useful to identify TDS contributions from gas-bearing shale samples compared to the coal-associated rocks.

Author contributions

Luis Castillo Meza – formal analysis, writing – original draft, writing – review and editing; Charles Cravotta – formal analysis, writing – original draft, writing – review and editing; Travis L. Tasker – formal analysis; Nathaniel Warner – writing – original draft, writing – review and editing; Lee Daniels – formal analysis, writing – review and editing; Zenah Orndorf – formal analysis, writing – review and editing; Tim Bergstresser – formal analysis; Amy Douglass – formal analysis; George Kimble – formal analysis; Joelle Streczywilk – formal analysis; Chris Barton – formal analysis, writing – review and editing; Stephanie Fulton – formal analysis, writing – review and editing; Aaron Thompson – formal analysis, writing – review and editing; and William D. Burgos – supervision, writing – original draft, writing – review and editing.

Declaration of competing interestCOI

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at https://doi.org/10.1016/j.apgeochem.2020.104540.

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